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(E74-10729) EXPERIMENTAL EVALUATION OF
ATMOSPHERIC EFFECTS ON RADIOMETRIC
MEASUREMENTS USING THE EREP OF SKYLAB
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EXPERIMENTAL EVALUATION OF ATMOSPHERIC EFFECTS
ON RADIOMETRIC MEASUREMENTS USING THE EREP
OF SKYLAB (EPN No. 439)
Contract No. NAS 9-13343

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EXPERIMENTAL EVALUATION OF ATMOSPHERIC EFFECTS ON RADIOMETRIC
MEASUREMENTS USING THE EREP OF SKYLAB

1. Program Summary

At the present time we are awaiting the results of the evaluation of our recently submitted proposal to extend the present contract.

2. Work Accomplished During the Reporting Period

In addition to the preparation of a proposal in response to a request from the Contracting Office (RTAO: BB631 B4), three other tasks were performed during the reporting period. These include:

- a. Completion of an "interim" final report to review and summarize the work completed for the year's period commencing 7 May 1974. This document was submitted along regular channels to the technical monitor.
- b. A comprehensive evaluation of the information received at the recently held SKYLAB EREP Principal Investigators Data Meeting with particular attention to how this study will be effected by problems encountered with the data. Conclusions drawn from this evaluation are summarized in Section 3.
- c. Study of the theoretical model adopted to compute atmospheric effect including a comparison of some preliminary calculations with S191 digital data. These results are included in Section 4.

3. Evaluation of Material From the PI Data Meeting

In a series of meetings and discussions with J. Barnes and J. Willard who attended the meeting and through study of the printed material received, a comprehensive evaluation of the impact of data/instrumental problems on fulfilling the scientific objectives of this study was undertaken. The results of this study are summarized by instrument.

3.1 S190 A Multispectral Camera - S190 B Earth Terrain Camera

Since data from these photographic instruments is utilized entirely to document the areal homogeneity of surface targets, anomalies such as streaks electrostatic marking, fogging, and emulsion lifting are of little consequence. These imperfections do not prevent imagery from being used to seek homogenous areas since for the most part, they are readily distinguishable against desert-like backgrounds.

3.2 S191 Interferometric Spectrometer

Data from the long wavelength (LWL) portion of the S191 Spectrometer scan was to have been used to deduce ground-truth data of atmospheric structure over the selected target sites. In addition, since the selected test site area was generally acquired up to 45° ahead of nadir, an effective viewing angle scan was performed by the astronaut as the site was kept within the instrument field of view providing a means of examining the angular isotropy of surface reflectance (to validate the Lambertian assumption made in the theoretical treatment) in the short wavelength (SWL) portion of the data. Examination of the instrumental error problems encountered leads us to believe that the LWL data may not be suitable for inferring vertical atmospheric structure. Absolute accuracies cited are ~15% in the SWL and thermal regions.

Of the problems cited in "S 191: Cautionary Note for Data Processed According to Photr 524 CH 2" forwarded to PI's recently, the most significant to our study is that of off band radiation. Atmospheric vertical structure (particularly the water vapor profile) should be obtainable from radiance measurements made in the 6-8 μm (6.7 μm H_2O Band) and 14-15 μm (15 μm CO_2 "hot" Bands). However, such inversion procedures are very sensitive to noise in the data. Since the offband radiation problem seems to become critical in precisely these absorption band spectral regions, it does not seem feasible to use the data for this purpose. The only recourse is to account for site atmospheric parameters using nearby conventional meteorological data. Climatology would probably provide better results than extremely noisy inversions. Due to cool down problems, there is no useful thermal data for Passes 2 or 5 which include both ground-truth test sites.

Data in the SWL region is affected by both calibration and off-band radiation problems, but appears to be useable in the region from .45 μm - 2.00 μm . Data from an acquired test site (Pass 43, Salton Sea Desert) for a period of ~3.7 seconds shows no apparent change in measured radiance with viewing angle. (See Figure 4.) However, these plotted data are not corrected for atmospheric effects.

3.3 S192 Multispectral Scanner

The low frequency and high frequency noise filtering techniques which have been developed are certainly more than adequate to satisfy the data requirements of this study. In selecting an areal segment of S192 digital data for analysis, areal averaging will most likely be done at a resolution of ~25-35 km over the homogeneous targets selected for study, this averaging will provide good characteristic numbers.

The only serious impact due to S192 anomalies involves the loss of data during Pass 2 due to improper spectral alignment and Pass 5 since C/D/P was not properly seated. Fortunately, alternate segments (Passes 39 and 43) were requested for these sites, but simultaneous ground truth has been lost.

4. Preliminary Theoretical Results

Theoretical values for the S192 channel radiances are calculable from the expression:

$$R^i(a,b,c) = \frac{1}{\pi} \int_{\lambda_1^i}^{\lambda_2^i} I(\lambda) \phi^i(\lambda) [T_\theta(\lambda) r(\lambda) \sin \theta (a T_Z(\lambda) - b) + c J_\theta(\lambda, r)] d\lambda$$

Radiances at zero air mass (such as are measured by aircraft ground truth) are similarly calculated by assigning unit transmissivity to the atmosphere on the exit path. The difference between the theoretical radiance and the zero air mass radiance is the atmospheric correction for the channel.
where: (Sources and references in brackets)

- 1) Superscript i refers to a particular S192 channel with wavelength limits specified by $(\lambda_1^i, \lambda_2^i)$: $i = 1, 2, 3, \dots, 13$.

2) Parameters (a,b,c) identify the calculated intensity R^i in the following manner:

$R^i(1,0,1)$ = theoretical S192 channel radiance

$R^i(0, -1, 0)$ = Theoretical "zero air mass" radiance

$R^i(1,1,1) = R^i(1,0,1) - R^i(0,-1,0)$ = atmospheric correction

3) $I(\lambda)$ = solar irradiance spectrum (watts $m^{-2} \mu m^{-1}$)
[Thekaekara, 1970]

4) $\phi^i(\lambda)$ = S192 i^{th} channel spectral responsivity
[NASA-JSC, 1973]

5) θ = solar elevation angle for the particular test site geometry

6) $T_\theta(\lambda)$ = atmospheric transmissivity from surface to space at angle θ

$T_Z(\lambda)$ = zenith atmospheric transmissivity
[Selby and McClatchey, 1972]

7) $r(\lambda)$ = surface reflectance spectrum
[ground truth and/or Hunt et al, 1970-1974]

8) $J_\theta(\lambda, r)$ = atmospheric diffuse reflectivity at angle θ and for surface reflectance r
[Fraser, 1973 and Plass and Kattawar, 1968]

These computations provide not only for immediate magnitude comparison with actual S192 radiances, but allow interchannel (differential) spectral effects to be evaluated.

Simulated S191 data may be obtained by evaluating only the integrand as a function of wavelength while assigning the response function an effective value of 1.0 at all wavelengths.

A sample calculation made with this scheme is illustrated in Figures 1, 2, 3. Figure 1 lists the input parameters used in the evaluation of the equation for radiance. The solar zenith angle is 37.55° . A summer mid-latitude atmospheric model was chosen with a 23 km visual range haze model. Atmospheric transmissivity (4 & 5) is calculated as a function of wave-number (1 & 2). Columns 3, 6, 7 are input, column 7 being carefully chosen to

compliment the atmospheric model and viewing geometry.

Unit band pass (i.e. square response function centered at given wave-number) quantities are given in Figure 2. Columns 1a, 1b, 1c are, respectively: theoretically satellite measured radiance, air craft (low flying) measured radiance, and atmosphereless radiance. This last quantity represents the radiance measured in the total absence of an atmosphere i.e., the reflectance spectrum of an identical surface area element illuminated by solar radiation in identical viewing geometry at the top of the atmosphere. Column 2 is the correction which must be subtracted from a satellite measurement to give the radiance which would be measured by an earth resources aircraft. Columns 3 b & c are the surface reflectance changes corresponding to the atmospheric effect over the zero air mass and atmospheric effect over the zero air mass and atmospheric cases, respectively. The larger this number is | Abs. Mag. | \leq 1, the greater the atmospheric effect. Some of these data (3b) are plotted in Figure 5. Note that the effect of Rayleigh scattering is significant, while the aerosols form a continuum. Column 4 would be input data of actual measurements if any [here (4)=(1a)] and Column 5 is the measured data corrected by Column 2 [in this case (5) = (1b)].

In Figure 3 these results are integrated over the S192 filter response functions which the program stores internally. The tabulated quantities are analogous to those in Figure 2 differing only in their bandpass weighting. The column labelled "GNUBAR" refers to the filter function centroid wavenumber as calculated from the stored filter function. Here again, the "MEASURED RADIANCE" column is a dummy set equal to the theoretical value since no data was available.

The results of these calculations (Figs. 2 and 3) are plotted in Figure 4 and compared to three time segments of S191 data from Pass #3 over the Salton Sea Desert. Note that the calculated values are too large by a factor of 10 (they are plotted as multiplied by 10^{-1}). The reason for this discrepancy is not known. The S191 data was taken from an Intermediate Radiance Parameters Tabulation (Product S042-3). The theoretical curve is somewhat low in the visible (.4-.7 μm), but this is undoubtedly due to the low visible reflection spectrum assumed for the surface (.20).

The column labelled "SURF. REF. CHANGE - ZERO AIR MASS" provides a measurement of the atmospheric correction which must be made in surface reflectance determinations to interpret space platform - based multi-spectral data in times of low altitude aircraft measurements. As expected the most significant atmospheric effect is in the visible and slopes down into the rear infrared. (Note that since only a few points are plotted, the effect of narrow absorption lines is not seen in the data.)

5. Significant Results

Since the calculations are of a preliminary nature there are no significant results to report at this time.

6. Further Plans

Further analysis will await receipt of requested data.

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FIGURE 1. INPUT PARAMETERS

SLANT PATH TO SPACE FROM ALTITUDE 0.0 KM ZENITH ANGLE = 37.55 DEGREES

MODEL ATMOSPHERE = 2 HAZE MODEL = 1 = 1.0 VISUAL RANGE

SOLAR ELEVATION ANGLE = 52.45 ASYMMETRY FACTOR = 1.00 DAY OF YEAR = 172, ORBITAL GEOMETRIC FACTOR = 1.033

WAVENUMBER	WAVELENGTH	SOLAR IRRADIANCE WATTS CM ⁻² UM ⁻¹ STR ⁻¹	ZENITH TRANSMISSIBILITY	TRANSMISSIVITY AT SOLAR ELEV. ANGLE	SURFACE REFLECTANCE	ATMOSPHERIC REFLECTANCE AT SOLAR ELEV. ANGLE
25000,	0.400	0.14764	0.5110	0.4287	0.120	0.071
24500,	0.408	0.17688	0.5279	0.4467	0.124	0.069
24000,	0.417	0.18235	0.5445	0.4644	0.128	0.067
23500,	0.426	0.17432	0.5607	0.4820	0.133	0.065
23000,	0.435	0.17171	0.5764	0.4990	0.138	0.063
22500,	0.444	0.19731	0.5916	0.5157	0.143	0.061
22000,	0.455	0.21205	0.6066	0.5322	0.148	0.059
21500,	0.465	0.21155	0.6205	0.5477	0.153	0.057
21000,	0.476	0.21193	0.6359	0.5626	0.159	0.055
20500,	0.488	0.20264	0.6463	0.5766	0.165	0.053
20000,	0.500	0.20064	0.6574	0.5891	0.171	0.051
19500,	0.513	0.19157	0.6681	0.6012	0.178	0.050
19000,	0.526	0.19107	0.6771	0.6114	0.184	0.048
18500,	0.541	0.18388	0.6851	0.6206	0.192	0.046
18000,	0.556	0.17742	0.6923	0.6289	0.199	0.045
17500,	0.571	0.17709	0.6980	0.6354	0.207	0.042
17000,	0.588	0.17584	0.7097	0.6488	0.216	0.040
16500,	0.606	0.16990	0.7169	0.6571	0.225	0.037
16000,	0.625	0.16385	0.7322	0.6749	0.235	0.035
15500,	0.645	0.15775	0.7486	0.6940	0.245	0.032
15000,	0.667	0.15145	0.7634	0.7114	0.256	0.030
14500,	0.690	0.14494	0.0001	0.0001	0.268	0.028
14000,	0.714	0.13752	0.0001	0.0001	0.280	0.028
13500,	0.741	0.12998	0.0001	0.0001	0.294	0.031
13000,	0.769	0.12264	0.0001	0.0001	0.308	0.034
12500,	0.800	0.11458	0.0001	0.0001	0.324	0.037
12000,	0.833	0.10624	0.7529	0.7060	0.341	0.042
11500,	0.870	0.09793	0.0001	0.0001	0.359	0.046
11000,	0.909	0.09102	0.6887	0.6423	0.150	0.023
10500,	0.952	0.08605	0.3343	0.2845	0.150	0.024
10000,	1.000	0.07728	0.8472	0.8113	0.150	0.025
9500,	1.053	0.06859	0.8562	0.8221	0.150	0.025
9000,	1.111	0.05989	0.5758	0.5231	0.150	0.026
8500,	1.176	0.05249	0.8253	0.7891	0.150	0.026
8000,	1.250	0.04525	0.8751	0.8450	0.150	0.027
7500,	1.333	0.03830	0.4418	0.3899	0.150	0.028
7000,	1.429	0.03332	0.0570	0.0365	0.150	0.028
6500,	1.538	0.02807	0.8905	0.8641	0.150	0.029
6000,	1.667	0.02231	0.8911	0.8651	0.150	0.030
5500,	1.818	0.01578	0.0021	0.0001	0.150	0.030
5000,	2.000	0.01064	0.2630	0.2167	0.150	0.031
4500,	2.222	0.00792	0.0001	0.0001	0.150	0.032
4000,	2.500	0.00568	0.0150	0.0098	0.150	0.032

FIGURE 2.

COMPUTATION OF UNIT BAND BASS RADIANCES
UNITS, RADIANCES = WATTS*CM⁻²*UM⁻¹ASTH⁻¹

WAVE- NUMBER	WAVE- LENGTH	THEORETICAL RADIANCE	ZERO AIR MASS RADIANCE	ATMOSPHERELESS RADIANCE	THEORETICAL CORRECTION	SURF,REF,CHANGE ZERO AIR MASS	SURF,REF,CHANGE NO ATMOSPHERE	MEASURED RADIANCE	CORRECTED RADIANCE
25000,	0.400	10	0.1155E-0	11b	0.1976E-02	11c	0.4610E-02	12	0.9575E-02
24500,	0.408	0.1360E-0	11b	0.2553E-02	0.5715E-02	0.1105E-01	0.5375E-00	3b	0.5813E-00
24000,	0.417	0.1379E-01	0.2832E-02	0.6097E-02	0.1096E-01	0.4973E-00	0.1621E-00	3c	0.1713E-00
23500,	0.426	0.1298E-01	0.2908E-02	0.6034E-02	0.1007E-01	0.4605E-00	0.1531E-00	4	0.1807E-00
23000,	0.435	0.1260E-01	0.3071E-02	0.6155E-02	0.9525E-02	0.4272E-00	0.1442E-00	4	0.1360E-01
22500,	0.444	0.1428E-01	0.3777E-02	0.7325E-02	0.1050E-01	0.3965E-00	0.1354E-00	5	0.2555E-02
22000,	0.455	0.1515E-01	0.4341E-02	0.8157E-02	0.1081E-01	0.3682E-00	0.1268E-00	4	0.1379E-01
21500,	0.465	0.1495E-01	0.4619E-02	0.8434E-02	0.1033E-01	0.3425E-00	0.1183E-00	4	0.2832E-02
21000,	0.476	0.1482E-01	0.4929E-02	0.8761E-02	0.9891E-02	0.3188E-00	0.1099E-00	4	0.1298E-01
20500,	0.488	0.1404E-01	0.5010E-02	0.8689E-02	0.9031E-02	0.2970E-00	0.1015E-00	4	0.2408E-02
20000,	0.500	0.1379E-01	0.5260E-02	0.8928E-02	0.8528E-02	0.2773E-00	0.9308E-01	4	0.5010E-02
19500,	0.513	0.1308E-01	0.5321E-02	0.8851E-02	0.7756E-02	0.2588E-00	0.8478E-01	4	0.5260E-02
19000,	0.526	0.1296E-01	0.5606E-02	0.9169E-02	0.7355E-02	0.2420E-00	0.7627E-01	4	0.5321E-02
18500,	0.541	0.1241E-01	0.5692E-02	0.9172E-02	0.6720E-02	0.2265E-00	0.6727E-01	4	0.5606E-02
18000,	0.556	0.1192E-01	0.5788E-02	0.9203E-02	0.6131E-02	0.2112E-00	0.5883E-01	4	0.5788E-02
17500,	0.571	0.1169E-01	0.6074E-02	0.9559E-02	0.5620E-02	0.1919E-00	0.4653E-01	4	0.6074E-02
17000,	0.588	0.1151E-01	0.6412E-02	0.9883E-02	0.5096E-02	0.1716E-00	0.3550E-01	4	0.6412E-02
16500,	0.606	0.1099E-01	0.6539E-02	0.9951E-02	0.4448E-02	0.1531E-00	0.2343E-01	4	0.6539E-02
16000,	0.625	0.1062E-01	0.6755E-02	0.1001E-01	0.3866E-02	0.1344E-00	0.1436E-01	4	0.6755E-02
15500,	0.645	0.1031E-01	0.6980E-02	0.1006E-01	0.3332E-02	0.1170E-00	0.6184E-02	4	0.6980E-02
15000,	0.667	0.1001E-01	0.7178E-02	0.1009E-01	0.2835E-02	0.1011E-00	0.1937E-02	4	0.7178E-02
14500,	0.690	0.4014E-02	0.1010E-05	0.1010E-01	0.4013E-02	0.1064E-04	0.1613E-00	4	0.4014E-02
14000,	0.714	0.3875E-02	0.1003E-05	0.1003E-01	0.3874E-02	0.1083E-04	0.1720E-00	4	0.1013E-05
13500,	0.741	0.3975E-02	0.9937E-06	0.9937E-02	0.3974E-02	0.1175E-04	0.1763E-00	4	0.3975E-02
13000,	0.769	0.4136E-02	0.9839E-06	0.9839E-02	0.4135E-02	0.1296E-04	0.1787E-00	4	0.4136E-02
12500,	0.800	0.4291E-02	0.9660E-06	0.9660E-02	0.4290E-02	0.1439E-04	0.1801E-00	4	0.9686E-06
12000,	0.833	0.9425E-02	0.6656E-02	0.9427E-02	0.2769E-02	0.1419E-00	0.8664E-04	4	0.9425E-02
11500,	0.870	0.4513E-02	0.9161E-06	0.9161E-02	0.4512E-02	0.1770E-04	0.1824E-00	4	0.9164E-06
11000,	0.909	0.3695E-02	0.2282E-02	0.3553E-02	0.1413E-02	0.9286E-01	0.5991E-02	4	0.2282E-02
10500,	0.952	0.2381E-02	0.9556E-03	0.3359E-02	0.1425E-02	0.2237E-00	0.4367E-01	4	0.9556E-03
10000,	1.000	0.3973E-02	0.2447E-02	0.3017E-02	0.1526E-02	0.9353E-01	0.4758E-01	4	0.2447E-02
9500,	1.053	0.3614E-02	0.2201E-02	0.2677E-02	0.1413E-02	0.9630E-01	0.5248E-01	4	0.2201E-02
9000,	1.111	0.2252E-02	0.1223E-02	0.2338E-02	0.1029E-02	0.1262E-00	0.5499E-02	4	0.1223E-02
8500,	1.176	0.2724E-02	0.1617E-02	0.2049E-02	0.1107E-02	0.1027E-00	0.4943E-01	4	0.1617E-02
8000,	1.250	0.2533E-02	0.1493E-02	0.1766E-02	0.1040E-02	0.1045E-00	0.6508E-01	4	0.1493E-02
7500,	1.333	0.1320E-02	0.5829E-03	0.1495E-02	0.7369E-03	0.1896E-00	0.1758E-01	4	0.1320E-02
7000,	1.429	0.9478E-03	0.4747E-04	0.1301E-02	0.9004E-03	0.2845E-01	0.4068E-01	4	0.4747E-04
6500,	1.538	0.1657E-02	0.9468E-03	0.1096E-02	0.7103E-03	0.1125E-00	0.7685E-01	4	0.9468E-03
6000,	1.667	0.1332E-02	0.7532E-03	0.8707E-03	0.5788E-03	0.1153E-00	0.7948E-01	4	0.7532E-03
5500,	1.818	0.4786E-03	0.6158E-07	0.6158E-03	0.4786E-03	0.1166E-04	0.3342E-01	4	0.6170E-07
5000,	2.000	0.5541E-03	0.9001E-04	0.4154E-03	0.2641E-03	0.4401E-00	0.2213E-01	4	0.9001E-04
4500,	2.222	0.2517E-03	0.3093E-07	0.3093E-03	0.2517E-03	0.1220E-04	0.2794E-01	4	0.3097E-07
4000,	2.500	0.1846E-03	0.2174E-05	0.2218E-03	0.1824E-03	0.1259E-02	0.2518E-01	4	0.2174E-05

FIGURE 3.

COMPUTATION OF 9192 CHANNEL WEIGHTED RADIANCES

CHAN NEL#	GNU- BAR	CHANNEL LIMITS	THEORETICAL RADIANCE	ZERO AIR MASS RADIANCE	ATMOSPHERELESS RADIANCE	THEORETICAL CORRECTION	SURF,REF,CHANGE ZERO AIR MASS	SURF,REF,CHANGE NO ATMOSPHERE	MEASURED RADIANCE	CORRECTED RADIANCE
1	23061,	20900, 25000,	0,1380E-01	0,3374E-02	0,1381E-01	0,1043E-01	0,4357E-00	0,1455E-00	0,1380E-01	0,3374E-02
2	21406,	18600, 23800,	0,1427E-01	0,4496E-02	0,1427E-01	0,9769E-02	0,3429E-00	0,1169E-00	0,1427E-01	0,4496E-02
3	19592,	17300, 21700,	0,1536E-01	0,5349E-02	0,1337E-01	0,8013E-02	0,2642E-00	0,8613E-01	0,1336E-01	0,5349E-02
4	18176,	15900, 20400,	0,1219E-01	0,5826E-02	0,1219E-01	0,6360E-02	0,2147E-00	0,6037E-01	0,1219E-01	0,5826E-02
5	16983,	15000, 18500,	0,1135E-01	0,6333E-02	0,1136E-01	0,5020E-02	0,1717E-00	0,3522E-01	0,1135E-01	0,6333E-02
6	15268,	15400, 16900,	0,9810E-02	0,6571E-02	0,9820E-02	0,3239E-02	0,6496E-02	0,7557E-02	0,9810E-02	0,6571E-02
7	13198,	10000, 15100,	0,7111E-02	0,3993E-02	0,7112E-02	0,3117E-02	0,5161E-03	0,7851E-01	0,7111E-02	0,3993E-02
8	10739,	9100, 12900,	0,4066E-02	0,1858E-02	0,4071E-02	0,2209E-02	0,4091E-03	0,3346E-01	0,4066E-02	0,1858E-02
9	9726,	7700, 11100,	0,3522E-02	0,1930E-02	0,3324E-02	0,1392E-02	0,1718E-02	0,2767E-01	0,3322E-02	0,1930E-02
10	8816,	7200, 10500,	0,2721E-02	0,1536E-02	0,2723E-02	0,1185E-02	0,1659E-02	0,3119E-01	0,2721E-02	0,1536E-02
11	6348,	5400, 7800,	0,1156E-02	0,4789E-03	0,1169E-02	0,6766E-03	0,3196E-03	0,2366E-01	0,1156E-02	0,4789E-03
12	4481,	4000, 5200,	0,3586E-03	0,1420E-03	0,3631E-03	0,2167E-03	0,3078E-03	0,2557E-01	0,3586E-03	0,1420E-03

Figure 4. RADIANCE VS. WAVELENGTH CALCULATIONS

THEORY: SUMMER MIDLATITUDE ATMOSPHERE
 PARAMETERS FOR PASS 43
 UNIT BAND PASS
 THEORETICAL S192 ①, ②, ③... (channel nos.)
 [ALL CALCULATIONS PLOTTED $\times 10^{-1}$]

DATA: S191 RADIANCES PASS 43 SALTON SEA - DESERT

SEGMENT 18:05:7.82-8.75

18:05:12.46-13.38

18:05:10.60-11.53



